Basin-scale overturning circulation and the role of small-scale processes (lessons learned from the Baltic Sea)

Lars Umlauf

Peter Holtermann, Hans Burchard, Toste Tanhua, Jim Ledwell, Oliver Schmale







Baltic Sea - Overview



- lateral scale: 1000 km
- depth scale: 100 m
- no tides
- connection to North Sea
- strong salinity gradients (0-20 g/kg)
- deep-water renewal by dense bottom currents
- inertial period ~14 hours
- 1st mode Rossby radius 3-10 km

Reissmann et al. (2009)

The role of mixing



Reissmann et al. (2009)

Internal-Wave Mixing (Bornholm Basin)



van der Lee and Umlauf (2011)

Late-Summer Conditions (September 2008)



Near-Inertial Motions



van der Lee and Umlauf (2011)

Summer Mixing (September 2008)



Winter Mixing (February 2010)



Scaling Internal-Wave Mixing



February 2010



September 2008

 $\varepsilon = \varepsilon_0 \frac{N}{N_0} \frac{S}{S_0} \qquad N_0 = S_0 = 3 \text{ cph}$

(van der Lee and Umlauf, 2011 MacKinnon and Gregg, 2003)

- Part of the shear is **sub-inertial**. Is this really an **internal-wave model**?
- Which motions cause the instability? Large-scale near-inertial waves or short internal waves?
- Is there an **internal-wave energy cascade** on the continental shelf?

Role of Boundary Mixing



- Observed interior mixing is too small to explain basin-scale mixing
- There is direct evidence for **boundary mixing**.
- So where does the mixing actually occur?

Deep-water mixing processes



- •Goal: quantify and understand deep-water mixing in the central Baltic Sea
- •Tracer experiment, turbulence measurements, moorings, modeling
- Mooring arrays and turbulence measurements
- High-resolution nested 3-D model (GETM)
 - 600 m lateral resolution
 - •vertically adaptive grid (Hofmeister et al. 2010)
 - •Second-moment turbulence closure model

Location and Hydrography



Lateral Tracer Spreading



Holtermann et al. (2012)

Vertical Mixing Rates

Initial stage Late stage (before boundary contact): (after boundary contact): $\kappa \sim 10^{-6} \text{ m}^2 \text{ s}^{-1}$ $\kappa \sim 10^{-5} \text{ m}^2 \text{ s}^{-1}$ $\frac{\partial c}{\partial t} = \frac{\partial}{\partial z} \left[\kappa_I \frac{\partial c}{\partial z} \right]$ $\frac{\partial c}{\partial t} = \frac{1}{A} \left[\frac{\partial}{\partial z} \left(A \kappa \frac{\partial c}{\partial z} \right) \right]$ Interior mixing 0 6 6 2 6 6 6 Intrusions **Boundary mixing**

Numerical Model Results



Seasonal Variability of Mixing



Holtermann and Umlauf (2012)

Deep-Water Dynamics



Holtermann and Umlauf (2012) Holtermann et al. (2014)

Sub-Inertial Variability



Holtermann et al. (2014)

Basin-Scale Topographic Waves





Conclusions



- •Near-inertial waves dominate internal-wave band
- Dissipation rates scale with N and S, similar to the shelf.
- Energetics and mechanisms of internal-wave mixing unclear
- Energy-consistent implementation unclear

